

1 **Picky Eaters: Carbon isotopic evidence for the uniform bioavailability of riverine**  
2 **dissolved organic matter to a model marine microorganism**

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14 **Key points**

- 15 • Riverine dissolved organic matter (DOM) from distinct rivers shows similar  
16 bioavailability to marine bacterium.  
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- 18 • Radiocarbon values of microbially respired carbon dioxide during incubations reveal  
19 preferential utilization of modern carbon compounds.  
20
- 21 • The fate of riverine DOM in coastal environments may depend on the metabolic  
22 potential of microorganisms that are present and active.

23

24 **Abstract**

25 Dissolved organic matter (DOM) is a key component of the global carbon cycle, with rivers  
26 delivering significant amounts of DOM to oceans. Urbanization and agricultural land-use alter  
27 the age and chemical composition of riverine DOM, which likely impact the downstream  
28 bioavailability of riverine DOM. Here, we use bioreactor incubations of a marine bacterium  
29 (*Pseudoalteromonas sp. 3D05*) to investigate DOM bioavailability from two distinct rivers: the  
30 Suwannee River (natural, non-urbanized), and the Upper Mississippi River Basin  
31 (anthropogenically influenced). We measured rates of microbial CO<sub>2</sub> production and radiocarbon  
32 ages (as  $\Delta^{14}\text{C}$ ) to assess bioavailable DOM remineralization. We observed nearly identical cell  
33 densities and degradation patterns for both riverine DOM incubations. Respired DOM  $\Delta^{14}\text{C}$   
34 values were also similar and decreased over time indicative of preferential utilization of recently  
35 synthesized “modern” substrates. These findings reveal unexpected similarities in riverine DOM  
36 bioavailability, indicating similar short term biological reactivity despite large DOM  
37 compositional differences.

38

## 39 **Plain Language Summary**

40 Our study explores the relationship between anthropogenic activity and breakdown of dissolved  
41 organic matter (DOM) from rivers, which represents a vital link between land and ocean  
42 ecosystems. The composition of DOM in rivers is linked to the characteristics of the surrounding  
43 land. Urbanization and agricultural land-use change the age and chemical composition of the  
44 riverine DOM. Consequently, these alterations induced from human activity would be expected  
45 to impact the bioavailability of riverine DOM to microorganisms in coastal environments. We  
46 compared DOM sourced from a natural river system (Suwannee River) to DOM from a river  
47 impacted by anthropogenic activity (Upper Mississippi River Basin) to understand how the  
48 availability of DOM from these distinct rivers varies to marine microorganisms. We carried out  
49 laboratory experiments with a model marine bacterium and measured the respiration of carbon  
50 dioxide and associated isotopic signatures during the breakdown of riverine DOM. Surprisingly,  
51 we discovered striking similarities in the breakdown patterns of DOM from both rivers, despite  
52 their differing origins. This suggests that the impact of human activities on downstream  
53 transformation of DOM may not be as straightforward as previously assumed and underscore the  
54 need for a nuanced understanding of how microorganisms process DOM in coastal  
55 environments.

## 56 **1. Introduction**

57 Rivers act as vital conduits for the transfer of organic carbon from terrestrial to marine  
58 ecosystems. Annually, it is estimated that substantial amounts of carbon (~250 Tg C) are  
59 transported to the global ocean via riverine dissolved organic matter (DOM) (Hedges et al.,  
60 1997; Schlesinger & Melack, 1981). Despite this significant input, a large proportion of this  
61 riverine material is remineralized by microorganisms and/or photochemically altered in the  
62 coastal ocean (Mopper et al., 2015; Moran et al., 2000). As a result, less than 10% of the DOM  
63 present in the ocean originates from terrestrial sources (Meyers-Schulte & Hedges, 1986; Opsahl  
64 & Benner, 1997; Williams & Druffel, 1987). Thus, the transformation of riverine DOM within  
65 coastal environments plays a crucial role in regulating the exchange of carbon and nutrient flows  
66 between terrestrial and marine ecosystems.

67 Riverine DOM is derived from leaf litter, grass, soil, freshwater algae as well as  
68 anthropogenic inputs. Consequently, the composition of riverine DOM is shaped by a myriad of  
69 factors including river morphology, basin lithology, nutrient loading, microbial activity, land-use  
70 and agriculture (human impacts; (Parr et al., 2015; Williams et al., 2010). In fact, human activity  
71 has been shown to severely impact the input, composition, and dynamics of riverine DOM (Parr  
72 et al., 2015; Vaughn et al., 2021; Williams et al., 2010). High resolution mass spectrometry has  
73 revealed that rivers originating from natural forest water sheds contain DOM with higher  
74 concentrations of condensed aromatics and polyphenolics which reflect greater contributions  
75 from soils and vegetation (Coppola et al., 2015, 2018; Opsahl & Benner, 1997; Vaughn et al.,  
76 2021). In contrast, rivers impacted by urbanization and agricultural activities exhibit different  
77 molecular signatures in the DOM, including a higher relative abundance of protein-like and  
78 aliphatic compounds as more sulfur- or nitrogen-containing heteroatoms (Roebuck Jr et al.,  
79 2019; Vaughn et al., 2021, 2023). Most riverine DOM is modern (e.g. soil organic carbon,  
80 vegetation debris) whereas anthropogenic inputs (e.g. wastewater, agrochemicals) and lithogenic

81 OC contributions (e.g. shales) are  $^{14}\text{C}$ -free which can lead to the re-introduction of aged carbon  
82 into the modern carbon cycle (Butman et al., 2012). Recent work suggests pre-aged DOM  
83 accounted for 3-9% of bulk riverine DOM exported from watersheds that were heavily impacted  
84 by human activity (Butman et al., 2015).

85 Historically, riverine DOM has been considered recalcitrant, or less bioavailable, for  
86 microbial consumption due to its origins from soils and aquatic plants (Blanchet et al., 2017;  
87 Moran & Hodson, 1990). Extensive research over the two decades has demonstrated the capacity  
88 of microorganisms, particularly heterotrophs, to consume and remineralize DOM (Amon &  
89 Benner, 1996; McCallister et al., 2004; Moran et al., 2000; Raymond & Bauer, 2001a, 2001b).  
90 The specific mechanisms that mediate the bioavailability of DOM to microorganisms is still  
91 unclear. A multitude of chemical and physical properties have been shown to influence the  
92 bioavailability of DOM in aquatic environments including chemical structure/diversity,  
93 molecular size and sample history (fresh vs. aged organic material). Previous studies have found  
94 that DOM influenced by urban and agricultural activities typically exhibits a lower molecular  
95 weight and a higher proportion of labile components such as aliphatic and peptide-like  
96 compounds (Parr et al., 2015; Wagner et al., 2015; Wilson & Xenopoulos, 2009). In contrast,  
97 DOM from pristine river environments tends to have a greater concentration of aromatic  
98 compounds such as polyphenols which are thought to be more resistant to microbial degradation  
99 (Butman et al., 2012; Wagner et al., 2015). Consequently, chemical alterations induced from  
100 human activity would be expected to impact the bioavailability and downstream transformation  
101 of riverine DOM in coastal areas (Riedel et al., 2016). To the best of our knowledge, this has yet  
102 to be directly tested via controlled microbial incubations studies.

103 In this study, we use a novel bioreactor system (Isotopic Carbon Respirometer-  
104 Bioreactor; IsoCaRB, (Beaupré et al., 2016) to investigate the bioavailability and carbon isotopic  
105 composition of respired riverine DOM to a model marine isolate. We focus on two contrasting

106 different river systems: the Suwannee River system which is a natural, unaltered river and the  
107 Upper Mississippi River system that is highly altered and anthropogenically impacted. We  
108 discuss these results in the framework of riverine DOM bioavailability, respiration rates and the  
109 isotopic composition of labile DOM removed during each incubation.

110

## 111 **2. Materials and Methods**

### 112 *2.1 Study locations and setting of riverine DOM*

113 DOM Samples from Suwannee River (SR) and the Upper Mississippi River (UMR) were  
114 sourced from the International Humic Substances Society (IHSS). The SR is a blackwater river  
115 that runs southwestward from the Okefenokee Swamp in Georgia to the Gulf of Mexico  
116 through Florida. This slow moving, river is characterized by high levels of terrestrially derived  
117 OM including high humic compound concentrations representative of low anthropogenic input  
118 (Cawley et al., 2013; Green et al., 2015). Suwannee River DOM (SR DOM; #2R101N) was  
119 collected in 2012 by the IHSS at the southernmost dam on the Suwannee River sill in Fargo,  
120 Georgia, a location chosen for its sparse human population (Figure 1; Green et al., 2015). SR  
121 DOM was extracted on-site with reverse osmosis systems, desalinated using cation exchange  
122 (CEX) and then subsequently, freeze-dried and homogenized for storage (Green et al., 2015). SR  
123 DOM underwent elemental analysis and was determined to contain ~50.7% C (IHSS; Table S1).

124 The Mississippi River is one of the largest river systems in the world, running southward  
125 and draining watersheds from about 30 states in the USA into the Gulf of Mexico. The Upper  
126 Mississippi River (UMR) system carved by glaciers thousands of years ago, is now highly  
127 altered by human actions converting the surrounding terrain into agricultural or urbanized land  
128 (Fremling et al., 1989; Vaughn et al., 2021). Since the 1950s, the UMR has experienced  
129 significant landcover transformations primarily due to the expansion of agricultural lands  
130 (Ramankutty & Foley, 1999; Schnitkey, 2013; Wright & Wimberly, 2013) and urbanization due

131 to a steady rise in population (Eathington, 2010). Upper Mississippi DOM (UMR DOM;  
132 #1R110N) was collected in 2013 in Minneapolis, Minnesota using combined reverse  
133 osmosis/electrodialysis system to process river water. Final concentrated samples were then  
134 desalinated and freeze dried. Elemental analysis indicates the UMR DOM contains ~49.98% C  
135 (IHSS). Compared to SR DOM, UMR DOM contains 1.85 times the nitrogen atoms and 1.47  
136 times the sulfur atoms, highlighting the anthropogenic input to the DOM of the Upper  
137 Mississippi River system (Table S1; IHSS; Wagner et al., 2015).

## 138 ***2.2 Cultivation of a model marine microbe***

139 DOM incubations were carried out with a bacterial isolate (*Pseudoalteromonas sp.* 3D05) that  
140 was previously isolated from coastal ocean water samples (Canoe Beach, Nahant, MA;  
141 42°25'11.5" N, 70°54'26.0" W) (Datta et al., 2016). This strain was selected based on its  
142 metabolic capability to degrade a wide spectrum of carbon substrates. Previous genomic analysis  
143 revealed that this marine isolate possesses multiple gene copies for extracellular enzymes  
144 involved in the hydrolysis of proteins, carbohydrates, and chitin (Mahmoudi et al., 2020). Cells  
145 were grown from frozen glycerol stocks that were thawed and added to 25 mL of marine broth  
146 2216 (Difco #279110) in a combusted 125 mL flask and left to shake (145 RPM) at room  
147 temperature until log-phase. Subsequently, 500  $\mu$ L of culture (1% inoculation) was transferred to  
148 50 mL of modified Tibbles-Rawling (T-R) minimal media (Table S2) with glucosamine (0.5%  
149 w/v) as a carbon source in a 250 mL flask and left to shake until log phase was reached.  
150 Subsequently, cells were transferred to fresh modified T-R minimal media and grown until  
151 reaching mid-log phase. Cell density was monitored by measuring optical density (OD) 600 nm,  
152 based on a calibration curve between OD and colony forming units (CFUs). Once cultures had  
153 reached the desired cell density, a total of 50 mL of cells were harvested for injection into the  
154 IsoCaRB system. Prior to injection, cells were washed two times with modified T-R media  
155 containing no carbon sources. Briefly, the culture was centrifuged for 10 minutes at 3000 xg

156 (Beckman Coulter Allegra X-30R Centrifuge) and the supernatant was decanted. The resulting  
157 cell pellet was resuspended in 1 mL modified T-R media containing no carbon source and  
158 injected into the IsoCaRB system using a sterile 3 mL syringe (BD Biosciences # 309657) and a  
159 20-gauge needle (BD PrecisionGlide™).

### 160 **2.3. Bioreactor incubations**

161 A series of incubations using SR DOM or UMR DOM was carried out using the IsoCaRB  
162 system. The IsoCaRB system is comprised of a gas delivery and purification system, a custom  
163 Pyrex culture vessel, an inline CO<sub>2</sub> detector and integrated LabVIEW data-logging program,  
164 custom CO<sub>2</sub> traps, and a vacuum extraction line. Microbially respired CO<sub>2</sub> is continuously  
165 collected as successive fractions in custom molecular sieve traps. Subsequently, CO<sub>2</sub> is  
166 recovered from the traps by baking (530°C for 30 min) under vacuum within 24 h of collection,  
167 then cryogenically purified, quantified, and stored in flame-sealed Pyrex tubes for isotopic  
168 analysis (see 2.4). Each experiment is allowed to proceed until CO<sub>2</sub> concentrations resume near-  
169 baseline values. Gaseous CO<sub>2</sub> concentration measurements are corrected for baseline drifts and  
170 then rescaled to agree with the higher-precision manometric yields obtained from the trapped  
171 CO<sub>2</sub> as previously described (Beaupré et al., 2016). The normalized CO<sub>2</sub> concentrations are then  
172 corrected for the confounding effects of mixing in the culture vessel headspace and decreasing  
173 media volume to constrain the rate of CO<sub>2</sub> generation per unit volume of growth medium ( $\mu\text{g C}$   
174  $\text{L}^{-1} \text{min}^{-1}$ ), which serves as a proxy for the microbial CO<sub>2</sub> production rate. Details regarding the  
175 standard operating procedure for the IsoCaRB system, including sterilization and assembly,  
176 sample preparation, and CO<sub>2</sub> collection and purification are described in Beaupré et al. (2016).

177 For each experiment, 700 mg of SR DOM or UMR DOM and live cells were incubated at  
178 room temperature (~22°C) in the IsoCaRB system. Based on the elemental composition of  
179 samples, this resulted in a total DOM concentration of 14.8 mM and 14.6 mM in the UMR DOM  
180 and SR DOM incubations, respectively. The slurry was continuously stirred (90 rpm) under

181 aerobic conditions to provide an unlimited supply of O<sub>2</sub>. A total of five CO<sub>2</sub> fractions for Δ<sup>14</sup>C  
182 analysis was collected during incubation with UMR DOM and SR DOM (Table S3). In addition,  
183 the slurry was subsampled via the sampling port every 12 h to track the number of viable cells  
184 via plate counts. Approximately 100 μL of slurry was serially diluted to 10<sup>-5</sup>. Subsequently, 100  
185 μL was plated in triplicate onto MB2216 agar plates and spread with rattler beads (Zymo  
186 #S1001). Plates were allowed to grow at room temperature (~22°C), and colony forming units  
187 (CFU) were counted after 48 h to determine cell density at the time of sampling.

188 Abiotic production of CO<sub>2</sub> from the DOM samples resulting from off-gassing was  
189 quantified by carrying out a control incubation for both SR DOM and URM DOM samples in the  
190 absence of live cells. Approximately 700 mg of DOM (SR or UMR) and 2 l of modified T-B  
191 media were incubated in the IsoCaRB system. The DOM-media slurry was sparged for 72 hours  
192 in a manner identical to the incubations with live cells. After the completion of sparging, a  
193 molecular sieve trap was attached to the system to collect any CO<sub>2</sub> for the same duration of the  
194 incubations with live cells (~ 4 days). The resultant CO<sub>2</sub> fraction was quantified manometrically  
195 and sent for Δ<sup>14</sup>C analysis. The Δ<sup>14</sup>C values of microbially respired CO<sub>2</sub> that were observed  
196 during incubation with SR or UMR DOM were subsequently corrected using this respective  
197 value as described in Mahmoudi et al. (2017).

#### 198 ***2.4 Isotopic analysis of microbially respired CO<sub>2</sub>***

199 CO<sub>2</sub> fractions collected during incubations were sent to Keck-Carbon Cycle Accelerated Mass  
200 Spectrometry (AMS) facility at the University of California, Irvine for Δ<sup>14</sup>C analysis. CO<sub>2</sub> was  
201 converted to graphite using the sealed-tube Zn method (Walker & Xu, 2019; Xu et al., 2007).  
202 Radiocarbon values (<sup>14</sup>C) are reported in Δ<sup>14</sup>C notation and corrected for year of measurement  
203 following the conventions set forth by Stuiver and Polach (1977). Prior to proceeding with the  
204 DOM incubations, *Pseudoalteromonas sp.* 3D05 was incubated in the IsoCaRB system with an  
205 isotopically characterized carbon substrate (glucosamine; Δ<sup>14</sup>C = +33‰ ± 1.3). This incubation

206 was done in an identical manner to the DOM incubations. Analysis of the resulting CO<sub>2</sub> fraction  
207 from this test was determined to be  $\Delta^{14}\text{C} = +30\text{‰} \pm 1.3$ , which confirmed the robustness of our  
208 experimental set up.

209

### 210 **3. Results and Discussion**

#### 211 *3.1 Microbial remineralization of DOM from distinct river systems*

212 Anthropogenic activity can alter DOM composition (Jaffé et al., 2012; Lambert et al., 2015;  
213 Mattsson et al., 2009; Riedel et al., 2016; Wagner et al., 2015), but exactly how this may impact  
214 downstream processing by microorganisms in coastal and marine environments has yet to be  
215 tested. We experimentally tested how the bioavailability and remineralization of riverine DOM  
216 may vary to marine microorganisms due to contrasting watershed characteristics. The UMR has  
217 been significantly altered by human activities, with its surrounding areas converted into  
218 agricultural and urbanized land (Ramankutty & Foley, 1999; Schnitkey, 2013; Wright &  
219 Wimberly, 2013). Conversely, the SR is characterized by high levels of terrestrially derived  
220 organic matter and humic compounds, reflecting minimal anthropogenic impact. Interestingly,  
221 we observed nearly identical respiration patterns during bioreactor incubations with  
222 *Pseudoalteromonas sp.* 3D05 and SR DOM or UMR DOM. CO<sub>2</sub> production quickly increased to  
223 peak 0.7-0.9  $\mu\text{g C L}^{-1} \text{ min}$  within the first 10-16 hours of each incubation and then, progressively  
224 decreased to near baseline CO<sub>2</sub> levels within 4.5 days. Similarly, cell densities were comparable  
225 between both SR DOM and UMR DOM incubations. There was a steady increase in cell density  
226 during the initial half of the incubation, reaching a peak of  $4.9\text{-}5.7 \times 10^7$  CFU/mL, followed by a  
227 gradual decline for the remainder of the incubation period (Fig. 2). This decline in cell density is  
228 consistent with previous bioreactor work (Mahmoudi et al., 2020) and expected for a batch  
229 system (i.e. no replenishment of nutrients).

230           The rate of CO<sub>2</sub> production by microorganisms reflects of balance between catabolic and  
231 anabolic processes – this is often referred to as carbon use efficiency (CUE). When  
232 microorganisms transform organic matter a fraction of carbon is incorporated into biomass and  
233 the rest is respired as CO<sub>2</sub>. The exact cellular mechanisms that underlie CUE remain unclear as it  
234 has been shown to vary with environmental factors (e.g. temperature), the supply and complexity  
235 of substrates (nutritional and elemental composition, as well as energy content), and on the  
236 biochemical pathway of degradation and assimilation (Keiblinger et al., 2010; Roels, 1980;  
237 Russell & Cook, 1995). For example, simple and easily metabolizable compounds may result in  
238 higher CUE as more assimilated carbon is directed toward biomass synthesis, leading to  
239 increased cell density, rather than CO<sub>2</sub> production. In contrast, the consumption of complex  
240 substrates that necessitates the production and secretion of extracellular enzymes would require  
241 more energy investment for degradation (Allison et al., 2014). This would result in lower CUE as  
242 a greater proportion of assimilated carbon is directed toward energy generation through  
243 respiration rather than biomass, resulting in a lower cell density. Our observations of nearly  
244 identical cell density and CO<sub>2</sub> respiration rates indicate that efficiency in which this bacterium  
245 assimilated carbon into biomass was similar irrespective of the pool of DOM. This implies that  
246 microorganisms accessed and consumed carbon substrates that were similar in their energy  
247 expenditure across both DOM pools.

248           The percent loss of DOM (based on the total quantity respired CO<sub>2</sub> from the total C pool;  
249 Table 1) showed striking similarity between incubations with SR and UMR DOM, with losses of  
250  $1.2 \pm 0.01\%$  and  $1.3 \pm 0.01\%$ , respectively. Previous microbial incubations involving samples  
251 from ten different rivers observed DOM losses of 21% ( $\pm 5\%$ ) over the span a year (Riedel et al.,  
252 2016). Given the shorter duration of our experiment (i.e. days), our observed DOM losses align  
253 with expectation for shorter incubation periods. For example, 28-day bottle incubations of  
254 Suwannee River samples resulted in a DOM loss of 3% (Textore et al., 2018). Nevertheless, the

255 similarity in total DOM loss between the SR and UMR DOM incubations is both surprising and  
256 unexpected. Considering that both incubations had the same experimental conditions (e.g.  
257 temperature, incubation time, genetically identical bacterial cells, cell density, nutrient  
258 conditions, etc.), DOM composition was assumed to be the only variable determining the  
259 riverine DOM bioavailability. Our findings suggest that the bioavailability of both DOM pools  
260 did not differ substantially. Past DOM chemical characterization studies have found that urban-  
261 and agriculturally-impacted DOM contains lower molecular weight, higher relative aliphatic and  
262 peptide-like compounds and lower aromatic compounds in comparison to pristine river  
263 environments (Parr et al., 2015; Vaughn et al., 2021; Wagner et al., 2015). For example, a recent  
264 study by Vaughn and co-workers (2023) conducted 28-day incubations using streams draining  
265 from three dominant landcovers (forested, agriculture and urban) in the Upper Mississippi River  
266 Basin. They observed greater DOM loss in riverine samples collected from urbanized ( $10 \pm$   
267  $4.4\%$ ) and/or agricultural ( $7.6 \pm 3.1\%$ ) landscapes compared to more pristine, forested  
268 landscapes ( $5.6 \pm 3.2\%$ ). If we assume anthropogenically influenced DOM is more bioavailable  
269 than natural riverine DOM (e.g. terrestrial plants and organic-rich soils) (Butman et al., 2012;  
270 D'Andrilli et al., 2015; Riedel et al., 2016; Textor et al., 2018; Wagner et al., 2015), then we  
271 would expect differences in the amount of DOM respired.

272 The bioavailability of DOM to microorganisms is dictated by complex interplay of  
273 chemical, physical, and biological factors that ultimately determine the rate and extent of its  
274 degradation and utilization in aquatic environments. Our results may imply two plausible  
275 interpretations. Firstly, they suggest a similar pool of bioavailable DOM—both in quantity and  
276 quality—across distinct river systems, which diverges from prior studies linking DOM reactivity  
277 to catchment properties (Butman et al., 2012; Riedel et al., 2016; Vaughn et al., 2023; Williams  
278 et al., 2010). Moreover, the reactive portion in DOM has been shown to be molecularly very  
279 dissimilar between rivers such that each system has a unique group of compounds that is labile or

280 produced (Riedel et al., 2016). Alternatively, our results suggest that the marine bacterium  
281 employed in our study might exhibit limitations in its capacity to degrade certain DOM  
282 substrates selectively. Microbes possess variable enzymatic capabilities that directly impacts  
283 their ability to break down specific molecular classes of DOM (Allison et al., 2014). Despite  
284 potential differences in the chemical composition of DOM between the two rivers, the marine  
285 bacterium in our incubations might be constrained to consuming similar compounds from both  
286 DOM pools. This aligns with the evolving hypothesis that the diversity and abundance of  
287 specific microbial species in a given environment impacts the degradation and remineralization  
288 of organic matter (Carlson et al., 2004; Glassman et al., 2018; Mahmoudi et al., 2020). While  
289 prior incubation studies have observed significant differences in DOM loss between natural and  
290 anthropogenically influenced river samples (Riedel et al., 2016; Vaughn et al., 2023), these  
291 studies employed the native river water as the inoculum which likely contain disparate microbial  
292 communities. As a result, the differences in DOM loss might be attributed to disparities in the  
293 enzymatic capabilities of the microbial populations in each river sample. Taken together, our  
294 results suggest that the remineralization of riverine DOM may be heavily influenced by the  
295 degradation pathways of the active microbial communities, rather than by variations in DOM  
296 composition originating from different watersheds.

### 297 *3.2 $\Delta^{14}\text{C}$ signatures of microbially respired $\text{CO}_2$*

298 Unraveling the relationships between the characteristics of DOM and its susceptibility to  
299 microbial utilization is complex and challenging to address. Natural abundance  $^{14}\text{C}$  analysis is a  
300 robust tool for resolving the sources and ages of natural organic matter consumed and  
301 assimilated by microorganisms (Mahmoudi et al., 2013a, 2013b; Mailloux et al., 2013; Mindorff  
302 et al., 2023; Pearson et al., 2008; Petsch et al., 2001; Wakeham et al., 2006). This approach is  
303 grounded in the fact that heterotrophic microorganisms carry the same  $^{14}\text{C}$  signatures as their  
304 carbon sources. Thus, the  $^{14}\text{C}$  signatures of both microbial cellular components (e.g., membrane

305 lipids) and respired CO<sub>2</sub> can be used to infer microbial utilization of isotopically-distinct carbon  
306 sources (Hayes, 2001). We employed this approach to evaluate the <sup>14</sup>C age of microbially  
307 respired CO<sub>2</sub> during our incubations thereby allowing us to directly compare the bioavailability  
308 of DOM to its <sup>14</sup>C age. The <sup>14</sup>C age of DOM is thought to be an important factor in terms how  
309 easily microorganisms can access and utilize the organic compounds. Older organic material is  
310 thought to have undergone various transformations that make it less accessible to  
311 microorganisms while younger material is composed of recently photosynthesized compounds  
312 that are more readily degradable compounds.

313         The observed  $\Delta^{14}\text{C}$  values of respired DOM CO<sub>2</sub> ranged from +34 to +14‰ for SR DOM  
314 and from +24 to +14‰ for UMR DOM. During both incubations, respired DOM  $\Delta^{14}\text{C}$  values  
315 became more negative over time – by precisely 10‰. Recent work suggests northern hemisphere  
316 CO<sub>2</sub>  $\Delta^{14}\text{C}$  values were between +30 and +40‰ in 2012 and 2013 when SR DOM and UMR  
317 DOM were collected (Hua et al., 2022). This indicates preferential utilization of recently  
318 synthesized “modern” riverine DOM by microorganisms in our incubations. These  $\Delta^{14}\text{C}$  findings  
319 are consistent with previous time series incubations (Raymond & Bauer, 2001c) and biomarker  
320 approaches (Cherrier et al., 1999) that have observed microorganisms preferentially consuming  
321 DOM that is younger than the bulk DOM pool during the degradation.

322         In our study, the observed  $\Delta^{14}\text{C}$  signatures of respired CO<sub>2</sub> are consistent with utilization  
323 of modern carbon compounds that are likely derived from recently photosynthesized sources  
324 (e.g. algae, plants, etc.). In the case of UMR DOM, these  $\Delta^{14}\text{C}$  values were very similar to the  
325 bulk DOM  $\Delta^{14}\text{C}$  value (29‰ ± 3), suggesting that microorganisms consumed the freshest  
326 material within this DOM pool. However, the  $\Delta^{14}\text{C}$  signatures of respired CO<sub>2</sub> during the SR  
327 DOM were more negative (by ~20‰) compared to the bulk DOM  $\Delta^{14}\text{C}$  value (+54 ± 3‰). This  
328 implies the presence of younger organic compounds, plant-derived compounds, in SR DOM that  
329 might be less accessible to *Pseudoalteromonas sp. 3D05*. SR receives substantial inputs of

330 terrestrially derived organic matter which is enriched lignin-derived polyphenols.  
331 *Pseudoalteromonas sp. 3D05* lacks genes for known enzymes (e.g. dioxygenases, xylosidase,  
332 cellobiase) that are needed to break down newly synthesized plant and terrestrial inputs. Thus, it  
333 is likely forced to consume compounds that may be slightly older in  $^{14}\text{C}$ -age (lower  $\Delta^{14}\text{C}$ ) but  
334 may be easier to access based on its enzymatic repertoire.

335

#### 336 **4. Summary and Implications**

337         The anticipated increase in precipitation and flood events with climate change may  
338 amplify the export of anthropogenically-influenced DOM exported to marine ecosystems  
339 (Bianchi et al., 2013; Parr et al., 2015). Our DOM bioreactor experiments, allow for the direct  
340 comparison of anthropogenically influenced vs. natural river DOM microbial remineralization in  
341 coastal ocean conditions. Our monoculture experiments reveal that the majority of terrestrial  
342 DOM can readily escape remineralization on short timescales, irrespective of variations in DOM  
343 composition stemming from different watersheds. Thus, our findings suggest that the  
344 remineralization of riverine DOM in coastal environments may largely depend on the metabolic  
345 potential of the microbial community is both present and active. This is in combination with  
346 environmental factors that might also facilitate access to labile DOM (e.g. photochemical  
347 oxidation, Moran & Zepp, 1997). In addition, our results further corroborate the concept of an  
348 age-reactivity continuum (e.g. Walker et al., 2016), wherein microorganisms initially consume  
349 the younger, more bioavailable compounds before progressing to metabolize progressively older  
350 DOM compounds. Future work examining the molecular level information (i.e., molecular-level  
351 formula, functional groups, and structures) of the DOM in our experiments will help shed light  
352 on mechanisms underlying this process. Subsequent studies could investigate bioavailability of  
353 riverine DOM using a broader range of model microorganisms, potentially working in synergy to  
354 enhance the transformation and removal of DOM.

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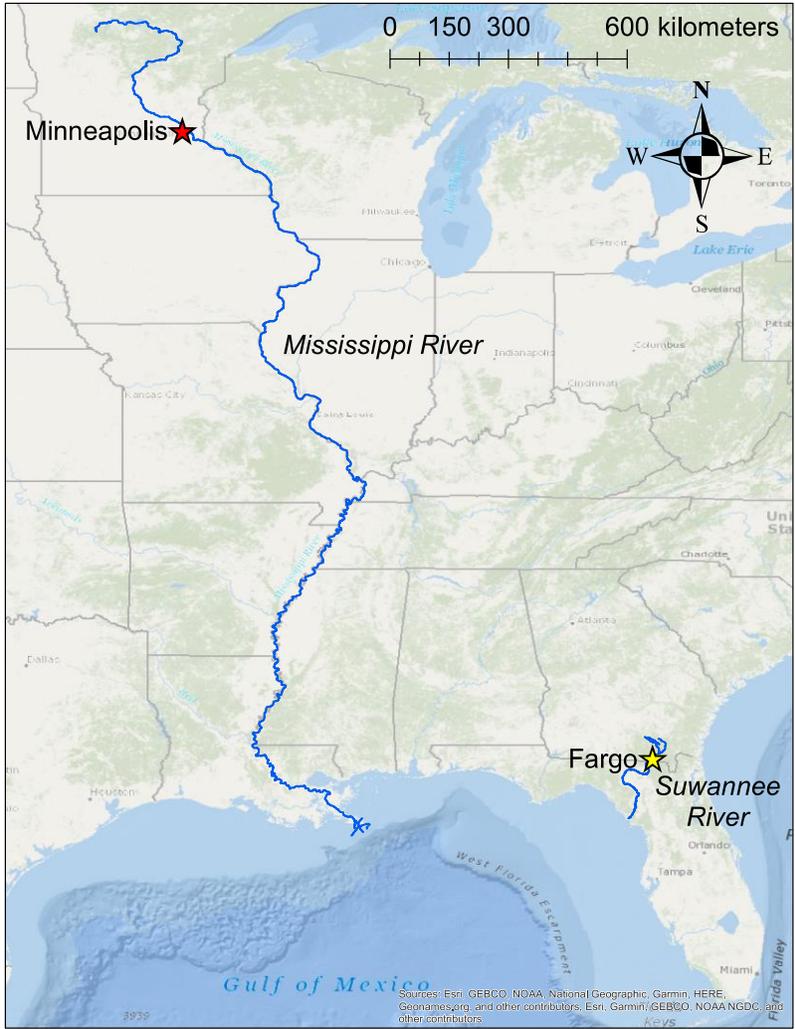
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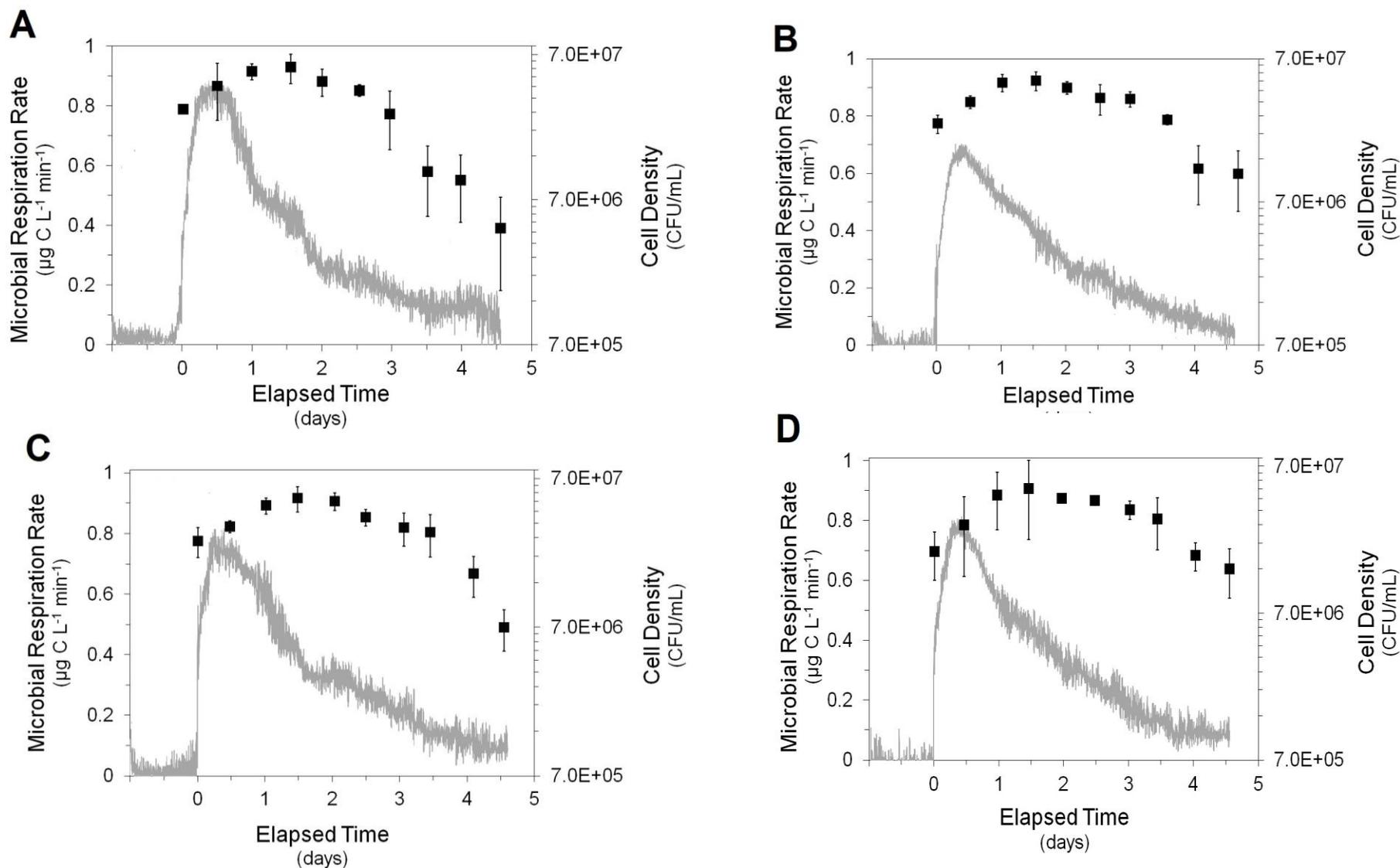
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364 **Open Research**

365 The genome for the bacterial isolate used in this study can be found on National Center for  
366 Biotechnology Information under BioProject ID PRJNA414740 under accession number  
367 PDUS00000000 (<https://www.ncbi.nlm.nih.gov/bioproject/?term=PRJNA414740>).



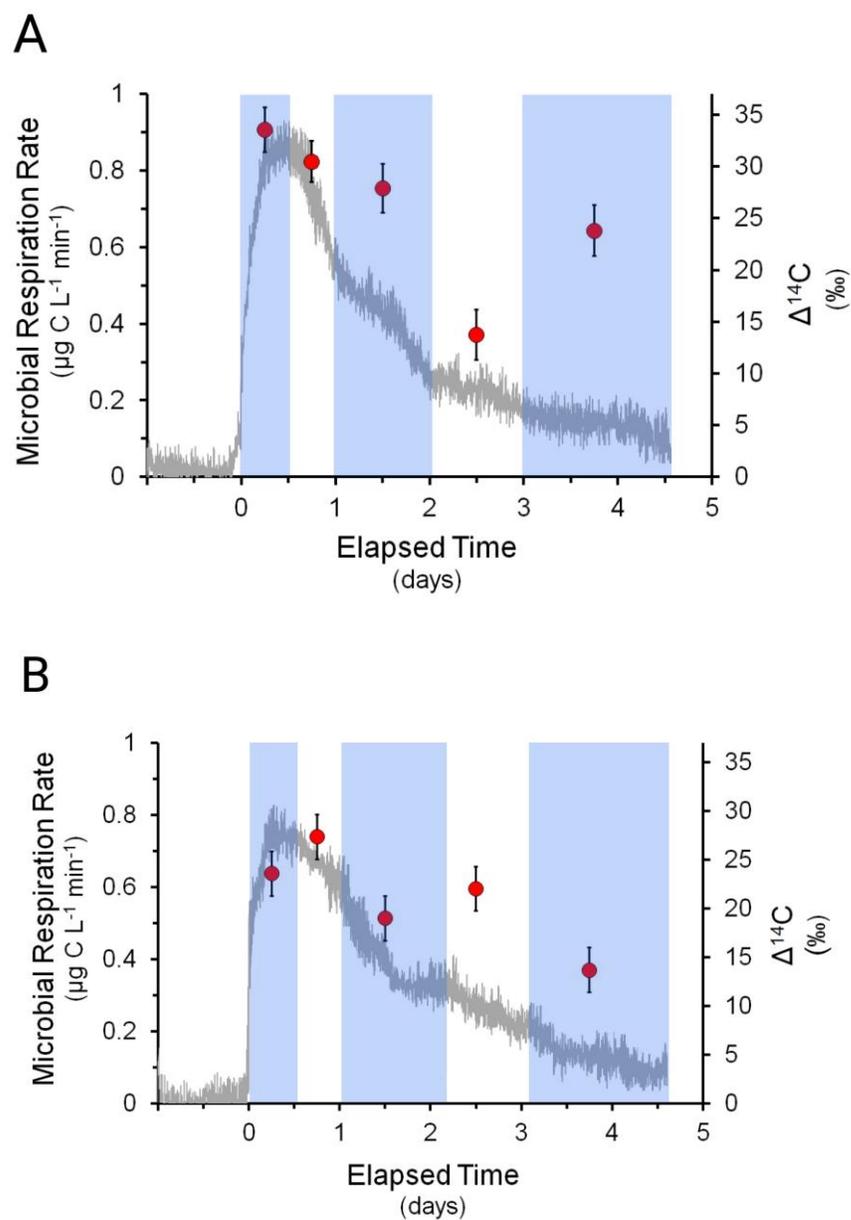
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 369 **Figure 1.** Map of the United States depicting location of rivers and collection sites for DOM.  
 370 Red star corresponds to the collection site for Upper Mississippi River DOM and yellow star  
 371 shows collection site for the Suwannee River DOM.



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374 **Figure 2.** Microbial respiration rates (grey line) and cell densities (squares) measured during replicate incubations of  
 375 *Pseudoalteromonas sp. 3DO5* with Suwannee River DOM incubation (A and B) and Upper Mississippi River DOM (C and D). Error  
 376 bars indicate standard deviations of the average (n=3).



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**Figure 3.** Microbial respiration rates (grey line) and  $\Delta^{14}\text{C}$  (red circles) signatures of respired  $\text{CO}_2$  observed during incubation of *Pseudoalteromonas sp. 3D05* with (A) Suwannee River DOM (SRDOM); and (B) Upper Mississippi River DOM (UMR DOM). The width of each blue box spans the time interval during which each  $\text{CO}_2$  fraction was collected for isotopic analysis, with the corresponding data point plotted at the mid-point for each fraction.

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